Understanding the environmental impacts of large fissure eruptions: Aerosol and gas emissions from the 2014–2015 Holuhraun eruption (Iceland)

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A R T I C L E   I N F O

Article history:
Received 21 March 2017
Received in revised form 14 May 2017
Accepted 17 May 2017
Available online 9 June 2017
Editor: F. Moynier

Keywords:
volcanic plume
volcanic emissions
volcanic eruption
air quality
environment
Iceland

A B S T R A C T

The 2014–2015 Holuhraun eruption in Iceland, emitted ~11 Tg of SO2 into the troposphere over 6 months, and caused one of the most intense and widespread volcanogenic air pollution events in centuries. This study provides a number of source terms for characterisation of plumes in large fissure eruptions, in Iceland and elsewhere. We characterised the chemistry of aerosol particle matter (PM) and gas in the Holuhraun plume, and its evolution as the plume dispersed, both via measurements and modelling. The plume was sampled at the eruptive vent, and in two populated areas in Iceland. The plume caused repeated air pollution events, exceeding hourly air quality standards (350 μg/m³) for SO2 on 88 occasions in Reykjahlíð town (100 km distance), and 34 occasions in Reykjavík capital area (250 km distance). Average daily concentration of volcanogenic PM sulphate exceeded 5 μg/m³ on 30 days in Reykjavík capital area, which is the maximum concentration measured during non-eruptive background interval. There are currently no established air quality standards for sulphate. Combining the results from direct sampling and dispersion modelling, we identified two types of plume impacting the downwind populated areas. The first type was characterised by high concentrations of both SO2 and 5-bearing PM, with a high SSO4/SPM (PM) (SO2/SPM > 10). The second type had a low SSO4/SPM ratio (<10). We suggest that this second type was a mature plume where sulphur had undergone significant gas-to-aerosol conversion in the atmosphere. Both types of plume were rich in fine aerosol (predominantly PM1 and PM2.5), sulphate (on average ~90% of the PM mass) and various trace species, including heavy metals. The fine size of the volcanic PM mass (75–80% in PM2.5), and the high environmental lability of its chemical components have potential adverse implications for environmental and health impacts. However, only the dispersion of volcanic SO2 was forecast in public warnings and operationally monitored during the eruption. We make a recommendation that sulphur gas-to-aerosol conversion processes, and a sufficiently large model domain to contain the transport of a tropospheric plume on the timescale of days be utilized for public health and environmental impact forecasting in future eruptions in Iceland and elsewhere in the world.

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1. Introduction

Large fissure eruptions (>1 km³) are one of the most environmentally hazardous volcanic scenarios due to their prodigious

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http://dx.doi.org/10.1016/j.epsl.2017.05.025
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release of gases and aerosol into the troposphere and lower stratosphere. This can impact air quality, terrestrial and aquatic environments, climate and health (Robock, 2000; Schmidt et al., 2011; Thordarson and Self, 2003; Witham et al., 2015). Icelandic large fissure eruptions were included in the UK National Risk Register in 2012 as one of the highest priority risks, due to the significant societal consequences caused by the Laki eruption 1783–1784 CE (15 km$^3$ lava, estimated 120 Tg SO$_2$, Thordarson and Self, 2003) across the northern hemisphere. During the Laki eruption, a thick sulphate aerosol veil and fluoride contamination of pasture and drinking water for livestock, led to the loss of ~20% of Iceland’s population through famine (Thordarson and Self, 2003). It has been suggested that mortality in the UK and mainland Europe also increased (Schmidt et al., 2011; Witham and Oppenheimer, 2004).

Large fissure eruptions are often assumed to be the closest analogues for large igneous province volcanism during the human era (e.g., Self et al., 2014) and so understanding them is further important in order to gain deeper insights into these massive environmental perturbations that punctuate Earth history. The Holuhraun eruption in Iceland in 2014–2015 (Fig. 1) is the largest such event since Laki and presents us with a rare opportunity to study volcanism of this style and scale using modern scientific techniques.

Tropospheric plumes from eruptions with low-level explosivity are complex mixtures of magmatic gases, entrained background air and aerosol particulate matter (PM), and usually low quantities of silicate ash. Volcanic plumes are known to contain nearly all naturally occurring elements (e.g., Allard et al., 2000; Symonds et al., 1992) ranging from the common metals (Na, K, Ca, Mg) to trace species (e.g., Pb, Mo, As, Se, Hg and Br). The interactions of this range of elements with ecosystems are highly variable and deposition of plumes can be either adverse or beneficial for terrestrial and aquatic environments (e.g. Cronin and Sharp, 2002; Jones and Gislason, 2008).

Volcanic gases and PM may also present a health hazard through air pollution. Sulphur dioxide (SO$_2$) levels in particular often exceed official air quality standards for this gas in the ambient air (Gislason et al., 2015; Schmidt et al., 2015; Tam et al., 2016). A large proportion of volcanic PM is in the PM$_{2.5}$ fraction (sub-2.5 μm diameter), which has become a key metric in air pollution guidelines. Short and long term exposure to PM$_{2.5}$ in the ambient air is estimated to cause over 3 million premature deaths globally per year (Lim et al., 2012), and it remains the single largest environmental health risk in Europe (Holgate, 2017). Detailed information on SO$_2$ and PM size-resolved mass flux, and gas and PM composition, is therefore necessary for reliable assessments of the effect of volcanic emissions on the environment and health.

The Holuhraun eruption occurred in the central highlands of Iceland (Fig. 1), on the ice-free part of Bárðarbunga volcanic system and lasted 6 months from 31 August 2014 to 28 February 2015. The estimates of its total SO$_2$ emission range between 11 Tg (Gislason et al., 2015) and 19 Tg (Gauthier et al., 2016), 1.5–2 times more than the total anthropogenic SO$_2$ emissions from the European Economic Area in 2011 (European Environmental Agency, 2017). Daily SO$_2$ gas fluxes averaged 60 kt/day (Gislason et al., 2015), a rate of outgassing rarely observed during sustained eruptions. It dwarfed most other long-term volcanic SO$_2$ emitters, such as Kilauea volcano in Hawaii (2–8 kt/d, Elias and Sutton, 2012) and Mt Etna, Italy (3.5 kt/d between 2005 and 2008, Salerno et al., 2009). It was one of the largest global tropospheric volcanic emissions since Laki, rivalled only by emission from Miyakejima 2000–2003 (~18 Mt SO$_2$, Kazahaya et al., 2004). This caused frequent and persistent deterioration of air quality in Iceland with regard to SO$_2$ (Gislason et al., 2015), and an impact on the chemistry of snow and rain precipitation around Iceland (Galeczka et al., 2016; Gislason et al., 2015; Stefansson et al., 2017). Long-range periodic impacts on air quality were reported in the UK, Ireland, Finland, the Netherlands and Austria (Gislason et al., 2015; Schmidt et al., 2015; Twigg et al., 2016).

Our study quantifies the impact of Holuhraun on populated areas in Iceland. Specifically we trace the evolution of the plume chemistry from the eruption site to 2 key areas of population: Reykjaehið, which is the nearest municipality to Holuhraun at 100 km distance, and Reykjavik capital area, which hosts ~60% of Iceland’s population, 250 km distance (Fig. 1). We present a novel dataset of gas and size-resolved aerosol chemistry, including a time series of constituents relevant for air quality assessment, i.e., SO$_2$, PM$_{2.5}$ and sulphate aerosol, and analysis of trace species including heavy metals. The findings of this study are relevant for multiple other countries impacted by volcanic emissions, either through prolonged eruptions, for example that of Kilauea in Hawaii, or through persistent quiescent degassing, such as from Masaya in Nicaragua.

2. Methodology

Data were collected at 3 sites (Fig. 1, section 2.1). The instrumentation and analytical techniques are detailed in sections 2.2–2.6.

2.1. Site descriptions

2.1.1. Holuhraun eruption site

The field campaign took place 21–22 January 2015, when we had favourable winter weather, dry and generally bright, thick
snow cover on the ground and −10 °C. The eruptive activity was exclusively effusive at the time, with lava issuing from the main vent Baugur and flowing towards N–NE (Fig. 1b). Due to the low wind speed (≤5 m/s) and high thermal flux of the eruption the plume grounded only periodically. We collected direct samples of the plume using filter packs (section 2.2) and cascade impactors (section 2.3), both on the ground in a dilute plume and in the highly concentrated airborne plume with the assistance of an Iceland coast guard helicopter (Table 1). We collected a repeat sample from a helicopter two months after the eruption finished to eliminate the possibility that the helicopter could be a source of detectable contamination.

2.1.2. Reykjahlíð town (100 km downwind)

Reykjhálið is a rural town with ∼300 inhabitants, located inland in north Iceland. It is the closest municipality to Holuhraun. Continuous automated monitoring of SO2 gas concentrations was operated by the Environment Agency of Iceland (EAI) both before and during the eruption (section 2.5). We obtained additional measurements during a field campaign 20–24 January 2015. Reykjahlíð was selected as the downwind location for the field campaign based on the plume-dispersion forecast issued by the Icelandic Meteorological Office (IMO). The plume was present at this location for most of the field campaign, as shown by the forecast and later supported by our data. Direct sampling of gas and aerosol (sections 2.2 and 2.3) and an optical particle counter (sections 2.4) were set up outside a local school, within ~100 m from the automated SO2 monitoring station.

2.1.3. Reykjavík capital area (250 km downwind)

Reykjavík capital area is located in south-west Iceland and has ~200,000 inhabitants, which is 60% of Iceland’s total population. It is a coastal city where non-volcanic air pollution is sourced mostly from anthropogenic activities (including one aluminium smelter) and sea-spray. Continuous automated monitoring of SO2 and PM2.5 was done both before and during the eruption at Hvalseyrarholt EAI station (section 2.5). Sulphate was analysed on filter samples collected daily at the same station (section 2.6).

### Table 1

<table>
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<tr>
<th>Location</th>
<th>Date</th>
<th>Instrument</th>
<th>Sample (S) or optical (O)</th>
<th>Sample ID</th>
<th>Distance (km) &amp; direction from main vent</th>
<th>Sampling duration (hh:mm)</th>
<th>Total volume (m³)</th>
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<td>Bulk aerosol and gas (S)</td>
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<td>FP</td>
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<td>Loc1_FP</td>
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<td></td>
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<td>20, N</td>
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<tr>
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<td>FP</td>
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</table>

FP = filter pack for bulk aerosol and gas sampling (section 2.2); SRK = 5 stage aerosol impactor; Moudi-13 = 13-stage impactor; Moudi-8 = 8-stage impactor (2.3); OPC = GRIMM optical particle counter (2.4); AQ = air quality station (2.5 and 2.6).

2.2. Filter packs

Filter packs were used to collect simultaneous samples of gas and PM. We followed the technique described by (Allen et al., 2000). The filter packs comprised one PM filter followed by 2–4 base-treated gas filters in an all-Teflon three-stage cartridge. The PM filter (Millipore, 47 mm, AAWP, pore size 0.8 µm) collects all particle size bins in bulk. The filters for gas sampling were Whatman 41 mm ashless circles impregnated with 5% K2CO3 and 1% glycerol in distilled deionized (DI) water, following the preparation technique described in Martin et al. (2010) and Pfeffer et al. (2013). This base treatment captures acidic gases (e.g., SO2, HF and HCl) on the filters. Blank filters were transported into the field and transferred to the sample bags in the same way as the sampling filters. This ensured that there was no unaccounted contamination due to transport and handling of the filters in the field. Airflow through the filter pack was generated using external pumps. After sampling, the filter pack was closed by a plastic lid to prevent particle loss and contamination. Once indoors, filters were then transferred into individual polypropylene bags (2 layers).

A three-stage sequential leaching scheme was employed to assess the solubility of metals in the PM, following the procedure used in Mather et al. (2012). In summary: (i) The first stage extracted the environmentally labile fraction by using ultraclean water for the leaching. (ii) The second stage leached moderately environmentally-labile material by adding a weak acid, hydroxylamine hydrochloride in 25% acetic acid to the residues from step 1. (iii) The third stage extracted the environmental immobile fraction of the PM (bound to silicates) by using HF. Following extraction, all samples were analysed for metals using an ICP-MS. Anions (Cl−, F− and SO42−) were analysed in the extracts from the first extraction stage only (water-soluble fraction) using ion chromatography (IC). The base-treated gas filters were analysed using IC only. All analysis was performed at the British Geological Survey laboratories in Keyworth, UK.

2.3. Aerosol cascade impactors

Cascade impactors allow size-segregated sampling of PM and have been used in multiple studies of volcanic aerosol (e.g., Ilyinskaya et al., 2012, 2010; Martin et al., 2011, 2010). A cascade
impactor consists of a series of stages with nozzles of decreasing diameter. An external pump attached to the impactor outlet creates an airflow through the stages. Beneath each stage is an impaction plate (where a filter is mounted) upon which the air jet impinges and where a fraction of the suspended aerosol particles is collected. Larger particles are deposited at the early separation stages whereas fine particles are not deposited until they have passed through correspondingly narrower nozzles in which the gas flows at higher velocities (e.g., Marple et al., 1991). The cut-off particle diameter for each impactor stage is conventionally reported as \(D_{50}\) which is the diameter of particles retained with 50% efficiency on each stage.

Here we used two types of cascade impactors (Table 1):

1. Micro Orifice Uniform Deposition Impactor (nano-MOI), model 12SNR, MSP corporation, operating at 10 l/min. It resolves 13 size bins between 18 and 0.01 μm cut-off diameter. PM larger than 18 μm is collected on the inlet stage, and PM <0.01 μm on the after-stage. Filters used were PTFE, 47 mm, 0.2 μm pore size. The last impaction stage (0.018–0.01 μm) was missing and not used during this campaign. This does not affect the PM cut-off sizes on the preceding stages.

2. MOUDI (model 130, MPS corporation), operating at 30 l/min. It resolves 8 size bins between 18 and 0.18 μm cut-off diameter. PM larger than 18 μm is collected on the inlet stage, and PM <0.18 μm on the after-stage. PTFE filters of 47 mm diameter and 1.0 μm pore size (Whatman 7590-004) were used for the main stages and a similar 37 mm filter (Whatman 7590-003) on the after-stage.

The filters (including procedural blanks) were handled and analysed using the same procedures as in section 2.2.

2.4. Optical particle counter

A GRIMM optical particle counter (model 1.107) was used in Reykjavík. The GRIMM data provides size segregated data for PM mass concentration and 8 particle size bins are resolved in the range 0.5 to >15 μm. Particle diameter data are automatically converted to particle volume using mean particle diameter between the thresholds of the 31 different channels, assuming spherical particles. Volume data are converted to a mass distribution using an instrument-specific density factor of 2.8 using standard protocols developed by GRIMM (Grimm and Eatough, 2009). Factory settings were used in this analysis.

2.5. Continuous SO2 and PM monitoring

Gaseous SO2 concentration is routinely monitored around Iceland by the EAI air quality stations. Automated SO2 stations both in Reykjavík and Reykjahlíð were in operation before the onset of the eruption. The SO2 measurements are done by an analyser equipped with a pulsed fluorescence spectroscopy meter, recording 10- and 60-min average concentration. All of the analysers are calibrated at 6-month intervals. The detection limit is 2 μg m\(^{-3}\) and the uncertainty is 5%.

PM\(_{10}\) and PM\(_{2.5}\) is automatically monitored at 14 stations in Iceland. Five of these are in the Reykjavík capital area. There is no permanent automated monitoring of PM at Reykjahlíð.

For Reykjavík, we analysed SO2, PM\(_{10}\) and PM\(_{2.5}\) data from Hvaleyrarholt station, which is also the only station to monitor sulphate through direct sampling (section 2.6). Non-volcanic background values of SO2, PM\(_{10}\) and PM\(_{2.5}\) were calculated based on the period from 2 January 2012 to 30 August 2014, an interval with no volcanic eruptions. New Year’s Eve (29 December to 1 January each year) was excluded from all calculations due to noticeable firework/bonfire pollution with PM\(_{2.5}\) reaching a maximum of 140 μg m\(^{-3}\) in 2012.

For Reykjahlíð, we analysed SO2 data from the automated station. Non-volcanic SO2 background was calculated based on the period from 2 January 2014 to 30 August 2014. Non-volcanic PM\(_{10}\) background was calculated from daily filter samples (24 h-average) collected in the same municipality between 2000 and 2005 (EAI data). We used only samples collected in the month of January as the PM mass concentrations fluctuate during the year, with highest concentrations during the summer months (May–August PM\(_{10}\) 7.3 ± 9.3 μg/m\(^3\)) and lowest during winter time (November–February PM\(_{10}\) 4.1 ± 3.7 μg/m\(^3\)) when snow cover minimizes re-suspension of rock dust, and agricultural and other anthropogenic activities (e.g. tourism traffic) are reduced. 1 January was excluded from the background concentrations due to New Year’s Eve pollution of up to 25 μg/m\(^3\) PM\(_{2.5}\). Background PM\(_{2.5}\) is not known but is assumed to be lower than Reykjavík (4.1 μg/m\(^3\)) as this is a rural inland area.

2.6. Sulphate monitoring (Reykjavík only)

Aerosol samples were collected using Digitel DHA-80 filter sampler located at Hvaleyrarholt station. The filters used (fibre glass and quartz) were not treated to collect gaseous phases. The filters were changed once every 24 h. After sampling, the filters were washed with DI water and peroxide solution and analysed for sulphate ion concentration using ion chromatography. The sample extraction and analysis were done at the Innovation Centre of Iceland. Non-volcanic sulphate background was calculated based on the period from 2 January 2014 to 30 August 2014.

3. Results

The chemical composition of the bulk plume (gas and aerosol phases) at the three sampling sites is summarised in Fig. 2. Full analytical results are in the supplementary material (Table SM and Fig. SM1).

3.1. Plume gas phase

The maximum gas concentrations were detected in the near-vent airborne plume, with 15-min average value of 43,000 μg/m\(^3\) SO2, 530 μg/m\(^3\) HCl and 210 μg/m\(^3\) HF. The average concentrations in the ground plume 4–20 km from the vent were 3,100 ± 2,900 μg/m\(^3\) SO2, 19 ± 2.3 μg/m\(^3\) HCl and 11 ± 5.9 μg/m\(^3\) HF.

In Reykjahlíð (100 km downwind) the hourly-mean and maximum concentrations of SO2 gas (26 ± 120 μg/m\(^3\), and 1700 μg/m\(^3\), respectively) during the eruption significantly exceed the pre-eruption background levels (0.41 ± 0.24 μg/m\(^3\), and 2.4 μg/m\(^3\), respectively), Fig. 3 and Table 2. During 20–24 January 2015, the gas composition in the ground plume was 370 ± 710 μg/m\(^3\) SO2, 4.5 ± 4.8 μg/m\(^3\) HCl and 1.5 ± 1.8 μg/m\(^3\) HF.

Reykjavík data (Fig. 3, 250 km downwind) were analysed as a function of the prevalent wind direction, as there is a noticeable direction-dependent difference in the background concentrations of gas and PM (Table 2). Background SO2 was higher during westerly winds compared to easterly winds (1.8 ± 1.4 vs 1.2 ± 0.8 μg/m\(^3\), respectively). The most likely sources of these pollutants from the west of the sampling location are the open sea (sea-spray) and an aluminium smelter. However, during the eruption, the increase in average SO2 is greater for the intervals of westerly winds than westerly winds, as the eruption site was located to the east of Reykjavík. The average and maximum concentrations of SO2 during the eruption (3.2 ± 0.8 and 15 μg/m\(^3\), respectively) significantly exceed the background levels. We considered the plume to be present in Reykjavík when total sulphur in the atmosphere (\(S_{\text{total}} = S_{\text{gas}} + S_{\text{PM}}\)) during a 24 h interval exceeded the maximum \(S_{\text{total}}\) measured before the eruption onset.
(30 μg/m³). Based on this analysis the volcanic plume was present in Reykjavik for 38 out of the 181 days of eruption. On those days, the average SO2 concentrations were 72 ± 140 μg/m³ (Table 2).

### 3.2. Plume aerosol phase

The near-vent airborne plume had extremely high mass concentrations of PM, with a 15 min average of 25,000 μg/m³. The sample collected in the near-vent airborne plume was not size-resolved, but the size-resolved plume sample collected in the near-vent grounding plume showed that PM$_{2.5}$ accounted for 75–83% of the total PM. PM$_{2.5}$ mass in the grounding plume at the eruption site was 60 μg/m³ (1 sample, concentration averaged 5 h).

The grounding downwind plume in Reykjavík town had an average PM$_{2.5}$ concentration of 13 ± 9.5 μg/m³ (70-h average value, Fig. 3). In Reykjavik capital area, the average daily PM$_{2.5}$ concentration throughout the whole eruption was 5.0±4.2 μg/m³, compared to 4.1 ± 3.6 before the eruption (Fig. 3). There was no significant difference between the daily and hourly averages of PM$_{2.5}$ concentration. The effect of the wind direction on the concentration of PM$_{2.5}$ was much less significant than on SO$_2$ (Table 2). The average sulphate concentration in Reykjavík during the eruption was twice as high as the pre-eruptive levels (3.0 ± 2.9 and 1.5 ± 1.1 μg/m³, respectively), with some wind field-dependent variations. During the 38 days when the plume was present in Reykjavik, SO$_4^{2-}$ accounted for 90% of the PM$_{2.5}$ mass on average, with the average SO$_4^{2-}$ concentration of 6.3 ± 3.7 μg/m³ (Table 2).

When interpreting the composition of highly dilute plumes sampled at more than a few km from source, such as at Reykjahlíð town, we have to account for a degree of mixing between the volcanic emissions and background air. In order to assess the extent of background air dilution of the Holuhraun plume at different sampling locations, we compare the sulphur to chlorine (S/Cl) molar ratio in volcanic PM with that of non-volcanic background PM, see supplementary material Fig. SM2. The PM collected near the emission source is concluded to have a predominantly volcanic chemical signature based on its S/Cl molar ratio (0.1–200, vs 0.05 in the background). The PM sampled in a more dilute plume 100 km downwind (Reykjahlíð) shows a higher degree of mixing with background PM (S/Cl 0.06–0.85), as is expected.

Different size fractions of the sampled PM appear to have different sources, volcanic and non-volcanic. The near-vent PM is of volcanic origin in all of the resolved size fractions (0.1 to >18 μm diameter), as shown by source appointment analysis described in Fig. SM2. The sample collected 100 km from eruption site (in Reykjahlíð) shows that PM < 1.8 μm has a predominantly volcanic composition, while PM 1.8–5.6 μm is mostly non-volcanic, although some degree of plume contribution cannot be excluded. PM ≥ 10 μm diameter shows a high degree of mixing between the two sources. This suggest that there is important surface controlled growth of the SO$_4^{2-}$ particles at relatively low supersaturation during long-range (>100 km) transport (Fig. SM2). Fig. 4 summarises the size-resolved chemistry of the aerosol phase, and its attributed...
source in samples from the eruption site and in Reykjahlíð. It shows that most of the analysed chemical species in the volcanic aerosol are in the PM₁₀ and PM₂.₅ size fractions.

4. Discussion

4.1. Plume composition

The gas phase mass ratios measured at the eruption site, HCl/SO₂ (0.026 ± 0.05) and HF/SO₂ (0.013 ± 0.02), are in good agreement with values (0.024 ± 0.07 and 0.011 ± 0.03, respectively) measured by Stefánsson et al. (2017), who also sampled the near-vent plume using filter packs. The standard deviation of gas concentrations between individual samples was generally high (Fig. 2), showing that the ground plume varied significantly in its gas phase concentration over short time scales. It is possible that the sampled gas was sourced partially from the degassing lava flows which generally has higher halogen/sulphur ratios (Thordarson et al., 1996), in addition to the gas from the main vent.

On average, the concentrations of the gas species decreased by ~1–2 orders of magnitude between the airborne plume and the grounding plume near the eruption site. The gas concentrations decreased by ~1 order of magnitude between sampling in the grounding plume near the eruption site and in Reykjahlíð at 100 km distance (Fig. 2).

By sampling the Holuhraun plume at varying distances from the eruptive vent we are able to show how sulphur in the volcanic plume fractionates between gas and aerosol phases over time. The concentration of the volcanogenic plume components decreases with time due to more extensive mixing and dilution by background air as the plume ages. The decrease in total S concentration (S_total = S_gas + S_PM) in the plume and the ensuing change in S_PM/S_gas ratio is used here as a proxy for an increasing plume age. Fig. 5 shows that S_PM/S_gas ratio is markedly higher when the plume is more dilute (i.e. older), with a strong negative correlation of ρ = −0.81. This is consistent with sulphur undergoing gas-to-particle conversion in the Holuhraun plume, reaching a S_PM/S_gas ratio of 0.62 in a sample taken at 100 km distance from the vent in Reykjahlíð 23–24 January 2015 (range of values 0.006–0.62), and up to 0.38 in Reykjavík at 250 km distance (range of values 0.016–0.38).

The relative abundance of sulphur and chlorine differs significantly between the gas and aerosol phases in the plume sampled shortly after its emission (Fig. 2). On average, the concentration of gaseous SO₂ exceeds HCl by a factor of 40. However, in the aerosol phase sulphur is less abundant than chlorine. It is known that S_PM/C_PM is controlled not just by the magmatic gas composition but also by the atmospheric conditions. Our previous work (Ilyinskaya et al., 2012) showed that on emission chloride condenses into the aerosol phase much more readily when the ambient temperature is low (<−5°C), while near-vent sulphate formation is unaffected, thereby lowering the S_PM/C_PM. This is applicable to the Holuhraun eruption as air temperatures were −5°C to −10°C at the time of sampling. Some mixing with background aerosol, which has a lower S_PM/C_PM ratio (0.05, Bruland, 1983) compared to the volcanic aerosol (0.37 ± 0.30, this study), is also occurring, as shown in the supplementary material (Fig. SM2).
Table 2
Impact of Holuhraun eruption on concentration of air pollutants (SO$_2$, SO$_4^{2-}$, and PM$_{2.5}$) in two populated areas. Reykjavík time series covers the whole eruption and is analysed here based on the prevalent wind direction, easterly (0–179°), and westerly (180–359°). ‘Plume present’ is identified when $S_{total}$ ($S_{gas}$ + $S_{PM}$) exceeded maximum background $S_{total}$. Reykjahlíð time series for PM$_{2.5}$ and SO$_4^{2-}$ covers 21–24 January 2015, during which the plume was present. Exceedance of air quality standards is indicated where data covered the whole eruption. While no air quality standards exist for SO$_4^{2-}$ we identified number of days exceeding daily-mean of 5 μg/m$^3$. Background data for PM$_{2.5}$ and SO$_4^{2-}$ does not exist for Reykjahlíð but they are considered to be lower than for Reykjavík due to its rural location. Units μg/m$^3$. ‘b.d.l.’ – below detection limits.

<table>
<thead>
<tr>
<th>Conditions</th>
<th>Reykjavík capital area, 250 km downwind</th>
<th>Reykjahlíð town, 100 km downwind</th>
</tr>
</thead>
<tbody>
<tr>
<td>Background (SO$<em>2$, PM$</em>{2.5}$ 02/01/2012–30/08/2014; SO$_4^{2-}$ 01/01/2014–30/08/2015)</td>
<td>Eruption (31/08/2014–28/02/2015)</td>
<td>Background (02/01/2014–30/08/2015)</td>
</tr>
<tr>
<td><strong>SO$_2$ (hour-mean)</strong></td>
<td><strong>Average σ Max Min</strong></td>
<td><strong>Average σ Max Min</strong></td>
</tr>
<tr>
<td>Easterly wind</td>
<td>0.8 1.4 38 b.d.l.</td>
<td>23 84 1400 b.d.l.</td>
</tr>
<tr>
<td>Westerly wind</td>
<td>1.9 4.4 59 b.d.l.</td>
<td>8.8 23 260 b.d.l.</td>
</tr>
<tr>
<td>Plume present (38 days)</td>
<td>n/a n/a n/a n/a</td>
<td>72 140 1400 b.d.l.</td>
</tr>
<tr>
<td>Air quality exceedance (1 h &gt; 350 μg/m$^3$ or 24 h &gt; 125 μg/m$^3$)</td>
<td>0 hours/7 days</td>
<td>34 hours/7 days</td>
</tr>
<tr>
<td><strong>PM$_{2.5}$ (hour-mean)</strong></td>
<td><strong>Average σ Max Min</strong></td>
<td><strong>Average σ Max Min</strong></td>
</tr>
<tr>
<td>Easterly wind</td>
<td>4.1 5.2 160 b.d.l.</td>
<td>4.8 5.2 91 b.d.l.</td>
</tr>
<tr>
<td>Westerly wind</td>
<td>4.0 5.1 160 b.d.l.</td>
<td>4.9 5.2 53 b.d.l.</td>
</tr>
<tr>
<td>Plume present (38 days)</td>
<td>n/a n/a n/a n/a</td>
<td>7.5 6.0 91 b.d.l.</td>
</tr>
<tr>
<td>Air quality exceedance (24 h &gt; 25 μg/m$^3$)</td>
<td>4 (in 32 months)</td>
<td>0</td>
</tr>
<tr>
<td><strong>SO$_4^{2-}$ (daily-mean)</strong></td>
<td><strong>Average σ Max Min</strong></td>
<td><strong>Average σ Max Min</strong></td>
</tr>
<tr>
<td>Easterly wind</td>
<td>1.2 0.8 3.5 0.2</td>
<td>3.2 3.0 15 0.0</td>
</tr>
<tr>
<td>Westerly wind</td>
<td>1.8 1.4 6.5 0.0</td>
<td>2.7 2.7 14 0.0</td>
</tr>
<tr>
<td>Plume present (38 days)</td>
<td>n/a n/a n/a n/a</td>
<td>6.3 3.7 15 0.0</td>
</tr>
<tr>
<td>&gt;5 μg/m$^3$ 8 days (in 8 months)</td>
<td>39 days, thereof 30 of volcanic source (in 6 months)</td>
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There are also differences in the size distribution of volcanic and background aerosol (Fig. 4b), as identified by the high time-resolution measurements from Reykjahlíð 21–24 January 2015. The plume was interpreted to be present during time intervals where PM$_{10}$ mass exceeded 10 μg/m$^3$. This is justified for this relatively short sampling interval during the time of year when non-volcanic aerosol is very low in this inland part of Iceland (background aerosol data presented in section 2.5), although this may not hold true for coastal areas which receive increased aerosol of marine origin during winter months (Gislason et al., 1996). Fig. 4 shows that the size distributions of PM mass concentration during plume presence is markedly different from the background. Within the plume the particles in the smallest aerosol size bins (〜0.5–3.2 μm diameter) are greatly enhanced compared to background. Furthermore, a coarse mode (〜7.5 μm diameter) is present in the plume and not in the background aerosol.

This shows that the finest and the coarsest PM size fractions (〜1.8 and 〜10 μm diameter) are predominantly volcanic. This is in good agreement with measurements of volcanic plumes from non-explosive activity elsewhere, in particular in locations characterised by sub-zero temperatures such as Erebus volcano in Antarctica (Ilyinskaya et al., 2010) and Fimmvörðuháls fissure eruption in Iceland (Ilyinskaya et al., 2012), where the dominant aerosol mode in a plume within a few minutes of emission is 〜0.2 μm in diameter. Plumes of similar age that are emitted into a warmer and more humid atmosphere have a dominant aerosol mode of 〜0.5–1.0 μm diameter, e.g. at Masaya in Nicaragua (Martin et al., 2011), Kiluaea in Hawaii (Mathet et al., 2012), and various volcanoes in Italy (Martin et al., 2008; Mathet et al., 2004).

Exposure standards are not established for most trace chemicals in airborne aerosol, although several trace species (V, Cr, Mn, Ni, As, Cd, Pb) have been tentatively linked to negative health impacts e.g. Chen and Lippmann (2009). Daily-mean concentrations for these trace species determined from samples collected in Reykjahlíð town over 4 consecutive days are shown in Fig. 6b. The species generally show little variation in concentration during the sampling period.

We assessed the environmental lability of the different species in the aerosol phase by a sequential leaching method. Fig. 6 shows the solubility of aerosol trace species. The figure also shows the enrichment factors (Mg-based EF, depicted by white circles on Fig. 6a) calculated based on the relative abundance of each element in the plume and in the magma, see e.g., Mather et al. (2012). Elements with EF>10 are considered as volatile. The majority of the measured trace species were soluble in water or weak acid. The exception were the lanthanides, which were all poorly soluble, and also showed a generally low volatility. These results imply a high environmental mobility of trace metals deposited as aerosol from the Holuhraun plume, with near-immediate dissolution. Poorly soluble compounds, such as the lanthanides, also dissolve but over a longer timescale (Oelkers and Gislasen, 2001).

4.2. Impact on air quality

Time series of major volcanogenic air pollutants, SO$_2$ gas, PM$_{2.5}$ and sulphate (SO$_2^{2-}$) aerosol were collected in two populated areas, Reykjahlíð town and Reykjavik capital area. PM$_{10}$ data are also available but we will focus our analysis on PM$_{2.5}$ which is the dominant size bin for volcanic aerosol as discussed in section 4.1. PM$_{2.5}$ is also a highly relevant factor impacting on human health, morbidity and mortality (Holgate, 2017).

4.2.1. Reykjavik capital area

During the course of the eruption, the European Commission (EC) air quality standards (EC Environment, 2016) for SO$_2$ exposure was exceeded 34 times for hourly limits (350 μg/m$^3$ hourly average) and 7 times for daily limits (125 μg/m$^3$ 24 h average), which also surpasses the number of permitted daily exceedances per year (max. 3 days). In the pre-eruption interval 2012–2014, SO$_2$ limits were never exceeded (Table 2).

Volcanogenic PM$_{2.5}$ mass concentrations in Reykjavik did not exceed the 24 hour-mean standard (25 μg/m$^3$) set by the World Health Organisation (WHO). The peak in PM$_{2.5}$ on 15 February 2015 (Fig. 7) is considered to be of a non-volcanic origin because...
Fig. 6. a) Concentrations of all major and trace elements in the aerosol phase as measured in the populated area of Reykjahlíð. The elements are shown in their solubility fractions*, normalised to 100%. Enrichment factor is also shown for elements where data exist. *Water soluble – milliQ extraction; Environmentally labile – weak acid extraction (HANCl); Silicate – HF extraction. b) Concentrations of trace metal pollutants which have been linked to negative health impacts in four samples collected in Reykjahlíð. The plot demonstrates that the difference in trace pollutant concentrations is relatively small between samples collected in a young plume (21–22 Jan. 2015) and mature plume (23–24 Jan. 2015).

Fig. 7. Time series of SO₂, PM₂.5 and sulphate (SO₄²⁻) in a) Reykjavík capital area during the whole eruption (24-h mean) and b) Reykjahlíð town 21–24 January 2015 (hourly-mean). Background sulphate concentration (mean + standard deviation) is shown by a grey bar (background data from Reykjavík used for both Reykjavík and Reykjalíð). Reykjalíð background is likely to be lower than Reykjavík due to its location.

it is not associated with an increase in sulphate and occurred after the eruption intensity had critically waned. The origin of this peak has not as yet been explained. It does not appear to be linked to a particular meteorological event (IMO, personal communication, February 2017). Potential explanations include nearby road or building works. The average hourly PM₂.5 concentration during the eruption (4.8–4.9 μg/m³) was elevated above pre-eruption background levels by 20% (Table 3). This is low compared to the percentage increase in SO₂ and sulphate aerosol (with increasing variation depending on wind direction, but approximately 2800% and 1700%, respectively). However, if we consider only the days when the plume was present in Reykjavík (n = 38, see section 3.1 for definition) the average hourly value of PM₂.5 is 7.5 μg/m³, which is 47% higher than the pre-eruption background.

The maximum hourly PM₂.5 concentration during the eruption period (91 μg/m³) was detected on one of the ‘plume-present’ days
(16 October 2014). But this PM$_{2.5}$ peak was lower than several peak concentrations during the non-eruptive period 2012–2014 (max 160 μg/m$^3$ per hour, Table 2).

However, emissions from the eruption caused a significant change in the chemical composition of the aerosol, with a large increase in sulphate mass, both in hourly-mean (∼170%) and peak concentrations (∼300%), Table 2. Unlike SO$_2$, air quality standards for sulphate are not available, and any definitions of sulphate ‘pollution events’ are somewhat arbitrary. In this discussion we set the limit at 5 μg/m$^3$, based on the suggestion of Lioy and Waldman (1989) in a review of anthropogenic pollution in major cities 1930s–1980s. Longo et al. (2005) used the same reference for identifying volcanic pollution zones around Kilauea volcano in Hawaii. Compared to the background aerosol in Reykjavík, 5 μg/m$^3$ is significantly higher than the background mean (1.5 μg/m$^3$, Table 2). During the 38 days when the plume was present in Reykjavík, SO$_4^{2-}$ accounted for 90% of the PM$_{2.5}$ mass on average; and 30 out of the 38 days experienced SO$_4^{2-}$ concentrations ≥5 μg/m$^3$ (compared to a total of 8 days during the pre-eruption period 2012–2014). The maximum SO$_4^{2-}$ daily average concentration (15 μg/m$^3$) was detected on 20 September 2014.

On multiple occasions throughout the eruption, PM$_{2.5}$ and sulphate mass concentration increased in Reykjavík without a corresponding increase in SO$_2$ mass concentration. For instance, there were at least 23 days when SO$_4^{2-}$ exceeded 5 μg/m$^3$ but SO$_2$ stayed within the EC standards (<125 μg/m$^3$). One of the best examples is 20 September 2014 (Fig. 7a). Based on the results from dispersion modelling presented in Schmidt et al. (2015), we show here (Fig. 8) that these events were caused by an arrival in Reykjavík capital area of a plume with a low S$_{gas}$/S$_{pm}$, where sulphur has near-complete gas-to-aerosol conversion. A plume with a low S$_{gas}$/S$_{pm}$ ratio is considered mature with respect to sulphur chemistry. On occasions, the ‘same’ plume inundated the same area twice, first as a primitive high S$_{gas}$/S$_{pm}$ plume and then several days later as a mature, low S$_{gas}$/S$_{pm}$, plume.

Importantly, the mature, low S$_{gas}$/S$_{pm}$ plume was not forecast as part of volcanic pollution monitoring during Holuhraun. Daily forecasts of the plume dispersion issued by the IMO accounted only for SO$_2$ concentrations in a primitive high S$_{gas}$/S$_{pm}$ plume. For example, the forecast for 20 September 2014 predicted good air quality (‘no plume’) in Reykjavík (http://brunnur.vedur.is/kort/calpuff/2014/09/19/00/calpuff_island_so2.html) but our results show that PM$_{2.5}$ exceeded the recommended levels, and sulphate was >5 μg/m$^3$ (Figs. 7a and 8b). Furthermore, results from dispersion modelling by (Schmidt et al., 2015) show an example of a ‘near-miss’ event (Fig. 8c), where a mature plume with concentrations of sulphate of 50–100 μg/m$^3$, narrowly passed the densely populated south-east coast of Iceland. Based on these findings, we recommend that mature plumes are included in air pollution forecasting during future eruptions.

4.2.2. Reykjahlíð town

The peak SO$_2$ concentrations and the frequency of air quality exceedance (hourly and daily means) were higher in Reykjahlíð (88 hours/10 days) than Reykjavík (34 hours/7 days, Table 3), as Reykjahlíð was located much closer to the eruption site. The PM$_{2.5}$ and sulphate time series are not as long-term in Reykjahlíð compared to Reykjavík. The data are nevertheless very important as they are a high resolution record of how the plume was fluctuating in intensity and composition at that location. The average and maximum hourly concentrations of both PM$_{2.5}$ and sulphate were 13 ± 9.5 and 11 ± 10 μg/m$^3$, respectively, which we consider to be significantly elevated above background (4 μg/m$^3$ and 1.5 μg/m$^3$, respectively, in Reykjavík but likely to be lower in Reykjahlíð as discussed in section 2.5). Daily-mean sulphate concentrations exceeded 5 μg/m$^3$ on 3 out of the 4 days of sampling (Fig. 7b).

In Reykjahlíð we made the same observation as in Reykjavík that the peaks in SO$_2$ and PM$_{2.5}$ were not always synchronous. The Reykjahlíð time line (Fig. 7b) shows two types of events: 1) SO$_2$ and PM$_{2.5}$ peak simultaneously at ∼5AM on 22 January 2015; and 2) PM$_{2.5}$ increases without a corresponding peak in SO$_2$ from ∼11AM on 23 January 2015. The source of the increase in PM$_{2.5}$ on 23 January is interpreted to be volcanic because sulphate is also elevated during the same time interval. This shows that there were episodes of a mature, low S$_{gas}$/S$_{pm}$ plume arriving in Reykjahlíð in the same way as was described in section 4.2.1 for Reykjavík capital area.

The concentrations of trace metals, including species that have been linked to negative health impacts, were similar when sampled in a primitive, high S$_{gas}$/S$_{pm}$ plume and a mature, low S$_{gas}$/S$_{pm}$ plume in Reykjahlíð (Fig. 6b). This suggests that mature Holuhraun plumes did not lose significant amounts of the metal-bearing aerosol, and were as potentially polluting as a more primitive plume with respect to these trace metals.

4.2.3. Implications for health impacts

Evidence suggests that the spread of tropospheric volcanogenic air pollutants during the Holuhraun eruption was reportedly the most extensive since the Laki eruption 1783–1784 CE. Holuhraun caused deterioration of air quality in populated areas >1000 km from source in Europe (Gíslason et al., 2015; Schmidt et al., 2015; Twigg et al., 2016). For comparison of impact intensity, SO$_2$ concentrations during an intense degassing period in 2008–2009 of Kilauea volcano, Hawaii, never exceeded the guidelines in the populated area of Kona in 80 km distance from the emission source. SO$_2$ guidelines have only been reported to have been exceeded within ∼30 km radius (Hawaii State Department of Health, Clean Air Branch, 2010).

While most studies on the relationship between volcanogenic pollution and health agree that the impacts are generally negative (e.g., Baxter et al., 1982, 1990; Longo et al., 2008; Tam et al., 2016), there is no good consensus on the exact cause, or the mechanisms involved. The published air quality exposure limits for pure SO$_2$ are designed based on laboratory studies in volunteers with non-severe asthma (US EPA National Center for Environmental Assessment, R.T.P.N., 2008). This has shown that asthma sufferers are particularly sensitive to low levels of SO$_2$, with symptoms including the rapid onset of constriction of the airways as is seen in asthma attacks. Air quality limits are set to protect this most vulnerable group. The evidence is weak on the long-term effects of repeated exposure to episodes of SO$_2$ pollution over time, such as occurs during prolonged eruptions. EC air quality standards are also not available for HCl or HF gas for periods over 8 h.

The health impacts of sulphate aerosol exposure in air pollution are much less understood. A major WHO review (WHO, 2013) concluded that there was no new toxicological evidence to support a causal role for inorganic sulphate in the health effects of PM$_{2.5}$, in accord with the comprehensive review on the subject by Schlesinger and Cassese (2003). However, epidemiological studies continue to report associations between sulphates and adverse impacts on human health. Postulated explanations include the possibility that sulphate is simply a marker for the causal agents in PM$_{2.5}$; or that toxic species may be formed in its chemical reactions with trace metals and the transition metals released in combustion processes. Sulphate has been reported to increase the solubility of iron in particles (Oakes et al., 2012), while metals may catalyse the production of sulphuric acid which in turn might react to produce toxicologically active soluble metal salts. The exact toxic mechanisms by which metals act are unknown, but free rad-
Fig. 8. Dispersion model of hourly-mean sulphur dioxide (SO$_2$, left) and sulphate aerosol (SO$_4^{2-}$, right) surface mass concentrations (μg/m$^3$) for selected days in September 2014, based on Schmidt et al. (2015) (full animation 31 Aug. 2014–28 Sept. 2014 in supplementary material SM3). Panels (a) and (b) show a mature low $S_{gas}/S_{PM}$ plume reaching the Reykjavik capital area on 14 and 20 September 2014 (right), whereas SO$_2$ mass concentrations are very low in the same location on those days (left). Panel (c) shows a ‘near-miss’ situation with a highly sulphate-rich mature plume narrowly passing the densely populated southeast coast of Iceland on 8 September 2014.

Volcanic plumes, as shown in this work and previous studies, are complex mixtures of gases and PM, some of which may be bio-reactive at low concentrations and add to the toxicity of the plume. The health impact of these, when combined in the ambient air, might be greater than expected from their individual properties, but this has not been studied – the possibility of gases and fine particles synergistically enhancing the toxicity of SO$_2$ in volcanic plumes cannot be ruled out.

The dispersion of a primitive plume with high $S_{gas}/S_{PM}$ ratio was operationally forecast during Holuhraun eruption. Anecdotal evidence suggests that the public heeded the forecasts and limited the time spent outdoors when the plume was forecast to be present (EAI and Icelandic Directorate of Health, pers comm, Jan-

[Diagram showing dispersion model of SO$_2$ and SO$_4^{2-}$]
uary 2015). However, our results show that on at least 18 days during the eruption the low $S_{\text{gas}}/S_{\text{PM}}$ mature plume was present in Reykjavik, while the forecast showed ‘no plume’. There is evidence to suggest that those days were noticed by local residents as having bad air quality. For example, on 20 September 2014 when sulphate concentrations in Reykjavik reached 15 $\mu$g/m$^3$, but $S_{\text{O}_2}$ was below exposure guidelines it was noted locally that: “This day was very remarkable, eyes and throat were burning but $S_{\text{O}_2}$...reached only 80 $\mu$g/m$^3$. Later during the eruption when $S_{\text{O}_2}$ exceeded 1000 $\mu$g/m$^3$ [in the same location] eyes were not burning as much as on September 20" (EAI, pers comm, August 2016).

Studies on the health impacts of the Holuhraun emissions should incorporate in their exposure assessments of at risk populations our findings on the variable concentrations of the $S_{\text{O}_2}$ and sulphate aerosol, e.g., comparing the health impacts of plumes with a high $S_{\text{gas}}/S_{\text{PM}}$ ratio (‘primitive plume’) and the low $S_{\text{gas}}/S_{\text{PM}}$ ratio (‘mature plume’).

5. Conclusions

Holuhraun was the largest ($1.6\,\text{km}^3$) and longest (6 months) volcanic eruption in Europe since the Laki eruption in late 18th century. The gas-and aerosol rich plume of Holuhraun reached populated areas of Iceland up to at least 250 km distance. Our measurements showed that both the airborne plume, and the plume sampled at ground level contained enhanced concentrations of major (e.g. $S_{\text{O}_2}$, sulphate), and trace species (e.g. heavy metals), which on many occasions, were elevated above background levels by several orders of magnitude. The majority of these species were soluble in water or weak acid, implying high environmental mobility. We conclude that Holuhraun eruption caused a change in the average chemical composition of the atmospheric aerosol that was detectable at least as far as the downwind location of the Reykjavik capital area (250 km), in particular during the first two months of the eruption. The volcanic aerosol size was predominantly in the fine ($\text{PM}_{2.5}$) size fraction with a significant portion in the ultrafine fraction ($\text{PM}_{1.0}$).

There were two types of volcanic plume detected in the populated areas that impacted air quality in different ways. The first type, a ‘primitive’, plume had high concentrations of both $S_{\text{O}_2}$ and sulphate, with a high $S_{\text{gas}}/S_{\text{PM}}$ ratio. The second type, a ‘mature’ plume had a low $S_{\text{gas}}/S_{\text{PM}}$ ratio, where sulphur had undergone significant gas-to-aerosol conversion. In the Reykjavik capital area (250 km downwind), there were at least 23 days when $S_{\text{O}_2}^2$ exceeded 5 $\mu$g/m$^3$, a concentration which previous studies have considered to constitute a pollution event (Lioy and Waldman, 1989; Longo et al., 2008) but $S_{\text{O}_2}$ stayed within the EC air quality standards ($\leq 125\,\mu$g/m$^3$). This was caused by an arrival of a mature plume. Dispersion modelling of the volcanicogenics pollutants shows that there were also near-miss events where a highly sulphate-rich mature plume narrowly passed the populated areas of Iceland without impacting upon air quality.

Both plume types had similar concentrations of trace metals (as measured in Reykjalýð town 100 km downwind), including species that have been linked to negative health effects of air pollution from anthropogenic sources. The conclusion is that both primitive and mature plumes have the potential to be a source of environmental and air pollution, and in future major eruptions in Iceland both types warrant further investigation as undertaken in this study.

An innovation in Holuhraun eruption was real-time plume modelling in Iceland which enabled the path of the primitive plume to be forecast in advance, and warnings issued to the population. We recommend that in any future large fissure eruptions consideration be given to extending the forecasting of air pollution events to include the low $S_{\text{gas}}/S_{\text{PM}}$ mature plumes, and careful assessment given to their possible wider dispersal.

As well as guiding in terms of future hazard and human health impact mitigation, this new understanding of the constituents of and chemical processing within the plumes from a large fissure eruptions has relevance in terms scaling up to Laki-scale events and the episodes of prolonged fissure activity (large igneous province volcanism) that punctuate the geological record.

Acknowledgements

This work was funded by the NERC urgency grant NE/M021130/1 ‘Source and longevity of sulphur in Icelandic flood basalt eruption plumes’. The Icelandic Coast Guard is most gratefully acknowledged for the vital assistance in the field. The authors thank Sue Loughlin, Ármann Hóskuldsson, Porsteinsson Jónsson, Ñjáll Fan- nar Reynisson, Richard Yeo, and the IES eruption team for assistance with organising and carrying out the field operations. Morten S. Riishuus is thanked for the photograph in Fig. 1. The research has also benefited from the funding of the European Research Council under the European Union’s Seventh Framework Programme (FP/2007–2013)/ERC grant agreement no. [279802]; European Community’s Seventh Framework Programme Grant No. 308377 (Project FUTUREVOLC), TM, BMCC, CO, ME and EI acknowledge support from the NERC-funded Centre for Observation and Modelling of Earthquakes, Volcanoes and Tectonics (COMET). PBJ is grateful to Professor Robert Maynard CBE for his advice on the respiratory health effects of sulphate aerosols. AS thanks the University of Birmingham for funding through the Elite Scholar scheme.

Appendix A. Supplementary material

Supplementary material related to this article can be found online at http://dx.doi.org/10.1016/j.epsl.2017.05.025.

References


Lippmann, R., Turner, 2014. Technological and epidemiological studies of cardiovascular effects of ambient air fine particulate matter (PM2.5) and its chemical compo-


